

PACIFIC EARTHQUAKE ENGINEERING RESEARCH CENTER

Nonlinear Horizontal Site Response for the NGA-West2 Project

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ABSTRACT

The nonlinear soil amplification models developed by Walling et al. (2008) are revisited for two main reasons: (a) the simulation database on which the models were developed has been updated and extended, with additional magnitudes and soil profiles, and (b) two alternatives for the input shaking parameter are explored – the Peak Ground Acceleration (PGA) and the spectral acceleration for each period, (Sa(T)). The benefits and limitations of each alternative are discussed.

The model is based on site amplification factors, relative to a V_{S30} =1170 m/sec site, computed for 53 base soil profiles using the RASCALS computer program. The base profiles include ten V_{S30} values (160 m/sec, 190 m/sec, 270 m/sec, 400 m/sec, 560 m/sec, 760 m/sec, 900 m/sec, 1170 m/sec, 2830 m/sec, and 3150 m/sec), up to eight soil depths (25 ft, 50 ft, 120 ft, 250 ft, 500 ft, 1000 ft, 2000 ft, 3000 ft), and four nonlinear soil models (EPRI and Peninsular range models for cohesionless soils and Imperial Valley and Bay Mud models for cohesive soils). For each soil profile, the site response is computed for 11 levels of input rock motion: PGA₁₁₇₀= 0.01g, 0.05g, 0.1g, 0.2g, 0.3g, 0.4g, 0.5g, 0.75g, 1.0, 1.25g, 1.5g, produced by three point-source magnitudes: M = 5.0, 6.0, 7.0.

For each base soil profile, the shear-wave velocities, layer thickness and the nonlinear soil properties (strain dependence of the G/G_{max} and hysteretic damping) are randomized. Correlation of the velocities and layer thickness variations are considered. For each base profile, source magnitude and input ground motion level, the median amplification and the standard deviation of the amplification are computed from 30 realizations of the soil profiles and material properties. These results are intended for use by the NGA developers to constrain the nonlinear scaling of the site response for the ground motion models.

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1 Introduction

The objective of the NGA project is to develop ground motion models for shallow crustal earthquakes in California that cover all relevant sources in California (excluding subduction earthquakes). In the past, the authors of the empirical ground motion models often set limits on their applicability based on the data set used to derive the models. However, these limits are typically ignored in the application of the models in PSHA because a PSHA needs to have the ground motion estimates for all relevant scenarios and for the site-specific site conditions.

Rather than having the hazard analyst blindly extrapolate the ground motion models, the NGA project required the developers of the models to extrapolate their models. To help the model developers in this extrapolation, analytical models were used to develop constraints on the ground motion scaling.

This report is divided into two main parts; Chapters 2 and 3 describe the results of a large suite of 1-D site response calculations that can be used by the NGA developers to constrain the nonlinear site amplification and to constrain the soil depth scaling for shallow soil sites. The objective of these chapters is to document the site amplification factors that were made available to the NGA developers. Chapter 4 presents and describes an update to the work conducted by Walling et al. (2008). The motivation for updating the horizontal nonlinear site amplification model is driven by two main reasons: (1) the simulation dataset was extended to include a wider range of possible conditions, and (2) the functional form is modified to include two variations – one with PGA and one with Sa(T) as the input shaking parameter.

2 Analytical Site Response Model and Input Parameters

2.1 ANALYTICAL MODEL

The site response calculations are conducted using the computer program RASCALS (Silva and Lee, 1987). The RASCALS program combines the stochastic point-source analytical model commonly used in seismology to define the source and path scaling of earthquakes with the equivalent-linear approach response commonly used in geotechnical engineering to estimate the non-linear behavior of soils. This method is summarized below.

2.1.1 Point-Source Model

The point-source model used in RASCALS is based on the omega-squared model with a constant stress-drop (e.g., Boore, 1983). Random vibration theory (RVT) is used to relate root-mean-square (RMS) values computed from the power spectrum to peak values of acceleration and oscillator response in the time domain (Silva and Lee, 1987).

The acceleration spectral density for outcrop rock, a(f), is given by:

$$a(f) = C \frac{f^2}{1 + (f/f_c)^2} \frac{M_o}{R} P(f) A(f) \exp\left(\frac{-\pi f R}{\beta_0 Q(f)}\right)$$
(2.1)

where f is frequency in Hz and

- Q(f) = frequency dependent quality factor (anelastic attenuation)
- A(f) = crustal amplification
- P(f) = high-frequency truncation filter
- $f_c = corner frequency (Hz)$

C is a constant which contains source region properties (density and shear-wave velocity) and accounts for the free-surface effect, the source radiation pattern, and the partition of the energy into two horizontal components:

$$C = \frac{1.1\pi}{\sqrt{2\rho_0 \beta_0^3}}$$
(2.2)

Source scaling is provided by specifying two parameters: the seismic moment (M_o) and the stress-drop ($\Delta\sigma$). The seismic moment is related to magnitude through the definition of moment magnitude, M:

$$\log M_0 = 1.5M + 16.05 \tag{2.3}$$

The stress-drop relates the corner frequency to the moment:

$$f_c = \beta_0 \left(\frac{\Delta \sigma}{8.44 M_0} \right) \tag{2.4}$$

The amplification accounts for the increase in wave amplitude due to the impedance contrast from the source to the rock outcrop. The amplification depends on the shear-wave velocity and density structure from the surface to the source depth.

The P(f) filter accounts for the observation that the high-frequency content decays faster than predicted by the omega-squared model. The Anderson and Hough (1984) model for the high frequency filter is adopted:

$$P(f) = \exp(-\pi kf) \tag{2.5}$$

The kappa can be related to the Q in the shallow crust:

$$k = \frac{H}{\overline{\beta}_R \overline{Q}_S}$$
(2.6)

where the over-bar indicates an average over a depth H.

2.1.2 Site Effects Model

To model soil and soft-rock response, an RVT-based equivalent-linear approach is used. The outcrop rock power spectral density from the point-source model described in Section 2.1.1 is propagated through a one-dimensional soil column. RVT is used to predict the peak time domain values of shear-strain based on the shear-strain power spectrum. In this sense, the procedure is analogous to the program SHAKE (Schnabel et al., 1972) except that in SHAKE, the peak shear-strains are measured in the time domain. The RVT approach does not use a time history, thereby eliminating the need to conduct multiple analyses using a suite of input time histories consistent with the outcrop rock power spectrum.

In addition to the shear-wave velocity profile, the site effects model requires models for the strain dependence of the normalized shear modulus (G/G_{max}) and the hysteretic damping. The sets of G/G_{max} and hysteretic damping models used in this study are described in Section 2.2.

2.1.3 Input Rock Motion

Three point-source earthquakes, with magnitudes 5.0, 6.0 and 7.0 and a constant stress drop of 50 bars, were used to define the reference rock outcrop motion. The distance for the point-source model was adjusted so that the PGA for the rock outcrop ($V_{S30}=1170$ m/sec profile described in Section 2.3) matched the desired PGA. For each combination of Magnitude and PGA_{Rock}, the point-source distance was kept fixed, and the velocity profile was replaced with the profile for the desired case. The kappa value for the input rock motion, however, was not kept fixed, as described below.

The kappa for recorded surface ground motion on generic soil and generic rock sites in California are similar with a kappa of about 0.04 sec. If the input rock motion kappa was held fixed at 0.04 sec, then for soil sites, the additional damping in the soil profile will lead to a larger kappa for the surface soil ground motions. Since the observed data does not show a larger kappa on California soil sites than on California rock sites, the kappa for the input rock motion (e.g., the kappa used in the point-source model) was reduced so that the total kappa for a 1000 ft profile would remain at 0.04 sec, consistent with observations. An illustration of this is shown in Figure 2.1.

For hard-rock sites, the kappa for surface recordings is lower than for soil sites, so this approach of maintaining the total kappa of 0.04 sec breaks down for hard-rock sites. Two sets of hard-rock site response calculations were computed: one with a kappa of 0.04 sec and one with a reduced kappa of 0.006 sec. The lower kappa value is consistent with the kappa for hard-rock sites in the eastern U.S.



Figure 2.1 Example of kappa for the point source model used for input rock motion.

2.2 G/GMAX AND HYSTERETIC DAMPING MODELS

Four sets of G/G_{max} and hysteretic damping models are used: two for cohesionless soils (EPRI and Peninsular Range) and two for cohesive soils (Imperial Valley and Bay Mud). These four models are described below.

2.2.1 Cohesionless Soil Models

The EPRI model was developed following the 1989 Loma Prieta earthquake using strong-motion data from three earthquakes: the 1979 Coyote Lake earthquake, the 1984 Morgan Hill earthquake, and the 1989 Loma Prieta earthquake. This model was validated by Silva et al. (1997) using strong motion recordings from 48 San Francisco Bay area cohesionless soil sites. The G/G_{max} and the hysteretic damping curves for the EPRI (1993) model are shown in Figure 2.2. The EPRI G/G_{max} and damping curves were developed for generic applications to cohesionless soils in the general range of gravelly sands to low plasticity silts or sandy clays.

Following the 1994 Northridge earthquake, the EPRI soil model was tested against the strong motion data from about 80 sites in the Los Angeles region that recorded the Northridge earthquake. This test showed that the EPRI model had greater nonlinearity than was observed in the Northridge data. As a result, a revised set of G/G_{max} and hysteretic damping curves was developed for cohesionless soils in the Peninsular Range based on the Northridge data. The G/G_{max} and the hysteretic damping curves for the Peninsular Range (PR) model are shown in Figure 2.3.



Figure 2.2 Modulus and damping curves for the EPRI soil model.



Figure 2.3 Modulus and damping curves for the PR soil model.

2.2.2 Cohesive Soil Models

For the Young Bay Muds and Old Bay Clay, the Vucetic and Dobry (1991) cohesive soil curves for a PI of 30% are used. The 30% value represents an average value for these cohesive soils. The G/Gmax and the hysteretic damping curves for the Young Bay Muds and Old Bay Clay (BM) are shown in Figure 2.4.

To cover the potential range in nonlinearity for cohesive soils, a second model based on the Imperial Valley data is used. Based on laboratory dynamic testing (Turner and Stokoe, 1982) and recordings from the 1979 Imperial Valley earthquake (Silva et al., 1997), these soils appear to behave much more linearly than the soft soils along the margins of the San Francisco Bay area and Mendocino, California (Silva et al., 1997). Although the Imperial Valley soils contain clays, typical PI values are less than 30% and the Vucetic and Dobry G/G_{max} and hysteretic damping curves appear to have too much nonlinearity to be consistent with the large peak acceleration values (about 0.5g) recorded during the 1979 Imperial Valley earthquake. As a result, a suite of G/G_{max} and hysteretic damping curves was developed for these soils based on the results of the Turner and Stoke (1982) laboratory dynamic testing and modeling of the ground motions (Silva et al., 1997). The G/G_{max} and the hysteretic damping curves for the Imperial Valley (IV) are shown in Figure 2.5. The IV curves remain nearly linear at high strain. Since neither laboratory testing nor the recorded ground motions resulted in strains exceeding about 0.1%, the curves are unconstrained at larger strain values. For high strain levels, the curves were linearly extrapolated.



Figure 2.4 Modulus and damping curves for the IV soil model.



Figure 2.5 Modulus and damping curves for the BM soil model.

2.3 SHEAR-WAVE VELOCITY PROFILES

The simulations were conducted for a range of soil profiles parameterized by the V_{S30} and the depth to $V_S=1000$ m/sec. The soil profiles are summarized in Table 2.1 and are plotted in Figure 2.6. The four models of the nonlinear material properties described in Section 2.2 were used for the soil profiles as indicated in Table 2.1. In all, 53 combinations of V_{S30} , soil depth, and nonlinear properties were evaluated. The 53 cases are listed in Table 2.2.

For each case listed in Table 2.2, the site response was conducted for three magnitudes: M5.0, M6.0, M7.0 and 11 different PGA values of the outcrop rock motion: 0.001g, 0.05g, 0.1g, 0.2g, 0.3g, 0.4g, 0.5g, 0.75g, 1.0g, 1.25g, 1.5g. For each run, the amplification with respect to V_{s30} =1170 m/sec was computed.

As noted in Section 2.1.3, the kappa for the input rock motion was adjusted to maintain a constant total kappa of 0.04 sec for deep soil profiles. The kappa values used for the input rock motion in each case are listed in Table 2.2. For hard-rock sites, two values of the total kappa were used: 0.04 sec and 0.006 sec. The lower total kappa reflects the observed reduction in total kappa for hard-rock sites.

V _{S30} (m/sec)	Soil Depth (ft)	Nonlinear Material Properties	Input Rock Kappa (sec)
160	25, 50, 120, 250, 500, 1000, 2000, 3000, 30-1000	ВМ	0.030
190	25, 50, 120, 250, 500, 1000, 2000, 3000, 30-1000	IV	0.032
190	30-1000	PR	0.032
270	25, 50, 120, 250, 500, 1000, 2000, 3000, 30-1000	500, 1000, 0-1000 PR	
270	30-1000	EPRI	0.034
400	25, 50,120, 250, 500, 1000, 30- 1000	PR	0.037
400	30-1000	EPRI	0.036
560	25, 50,120, 250, 500, 30-1000	PR	0.037
560	30-1000	EPRI	0.036
760	25, 50, 260	PR	0.038
760	260	EPRI	0.038
900	25, 260	PR	0.038
1170	0	Linear	0.040
3150	0	Linear	0.040
2830	0	Linear	0.006

Table 2.1Summary of site response cases.

Case	Depth to V _s =1.0 km/sec (ft)	V _{s30} (m/sec)	Soil Model	Input Rock Kappa(sec)
1	25	160	BM	0.030
2	50	160	BM	0.030
3	120	160	BM	0.030
4	250	160	BM	0.030
5	500	160	BM	0.030
6	1000	160	BM	0.030
7	2000	160	BM	0.030
8	3000	160	BM	0.030
9	30-1000	160	BM	0.030
10	25	190	IV	0.032
11	50	190	IV	0.032
12	120	190	IV	0.032
13	250	190	IV	0.032
14	500	190	IV	0.032
15	1000	190	IV	0.032
16	2000	190	IV	0.032
17	3000	190	IV	0.032
18	30-1000	190	IV	0.032
19	30-1000	190	PR	0.032
20	25	270	PR	0.035
21	50	270	PR	0.035
22	120	270	PR	0.035
23	250	270	PK	0.035
24	500	270	PK	0.035
25	2000	270	PK	0.035
20	2000	270	PK	0.035
27	20,1000	270	PR	0.035
20	30-1000	270	FDDI	0.033
29	25	400	DD	0.034
31	50	400	PR	0.037
32	120	400	PR	0.037
33	250	400	PR	0.037
34	500	400	PR	0.037
35	1000	400	PR	0.037
36	30-1000	400	PR	0.037
37	30-1000	400	EPRI	0.036
38	25	560	PR	0.037
39	50	560	PR	0.037
40	120	560	PR	0.037
41	250	560	PR	0.037
42	500	560	PR	0.037
43	30-1000	560	PR	0.037
44	30-1000	560	EPRI	0.036
45	25	760	PR	0.038
46	50	760	PR	0.038
47	260	760	PR	0.038
48	260	760	EPRI	0.038
49	25	900	PR	0.038
50	260	900	PR	0.038
51	20	1170	Linear	0.04
52	0	2830	Linear	0.006
53	0	3150	Linear	0.04

Table 2.2List of site response cases.





Figure 2.6 Shear-wave velocity profiles for the eight V_{S30} cases.

2.4 SOIL PROFILE RANDOMIZATION

Both the shear-wave velocities and the layer thicknesses are varied using correlation models based on an analysis of 557 measured shear-wave velocity profiles (Silva et al, 1997). The algorithm starts with a given base-case profile and generates a suite of random profiles about the base-case profile accounting for the correlations. The details of the procedure are given in Silva et al. (1997). The correlation models are summarized below.

2.4.1 Velocity Profile Randomization

The development of the randomization of the shear-wave velocity profile is described in Appendix C of Silva et al. (1996). The model is summarized here. For each base profile, the depth to the V_s =1.0 km/sec layer is randomized within plus and minus 50% of the depth for the base profile. Within the soil profile, the layer thicknesses are assumed to follow a Poisson process with a depth-dependent rate. The depth-dependent rate is modeled by

$$\lambda(h) = 1.98[h + 10.86]^{-0.89}$$
(2.7)

where λ is the rate of layer boundaries (m⁻¹) and h denotes the layer depth in m.

The layer velocities are developed by defining the normalized quantity, Z_i, given by

$$Z_{i} = \frac{\ln(V_{i}) - \ln(V_{median}(h_{i}))}{\sigma_{\ln V}}$$
(2.8)

The lognormal distribution of the velocities and the correlation among layers is modeled by a first-order auto-regressive model:

$$Z_{1} = \varepsilon_{1}$$

$$Z_{i} = \rho(h,t)Z_{i} - 1 + \sqrt{1 - \rho(h,t)^{2}} \varepsilon_{i} \quad for i > 1$$
(2.9)

where ρ is the serial auto-correlation coefficient of Z, and ϵ_i are independent normal variates with zero mean and unit variance.

$$\rho(h,t) = (1 - \rho_d(h))\rho_t(t) + \rho_d(h)$$
(2.10)

where ρ_d is the depth-dependent correlation and ρ_t is the thickness-dependent correlation and are given by

$$\rho_d(h) = \begin{cases} \rho_{200} \left(\frac{h + h_0}{200 + h_0} \right)^b & \text{for } h \le 200 \\ \\ \rho_{200} & \text{for } h > 200 \end{cases}$$
(2.11)

and

$$\rho_t(t) = \rho_0 \exp\left(\frac{-t}{\Delta}\right) \tag{2.12}$$

The model coefficients, σ_{lnV} , ρ_{200} , h_0 , b, ρ_0 , and Δ are listed in Table 2.3.

Deveneter	VS30 (m/sec) Range				
Parameter	> 750 m/sec	360 to 750	180 to 360	<180	
σ_{lnV}	0.36	0.27	0.31	0.37	
ρ₀	0.95	0.97	0.99	0.00	
Δ	3.4	3.8	3.9	5.0	
ρ ₂₀₀	0.42	1.00	0.98	0.50	
h ₀	0.0	0.0	0.0	0.0	
b	0.063	0.293	0.344	0.744	

 Table 2.3
 Coefficients for velocity profile randomization model.

2.4.2 Nonlinear Soil Property Randomization

The modulus reduction and damping curves were independently randomized about the base case values. A truncated log normal distribution was assumed with a standard deviation of 0.35 natural log units at a cyclic shear strain of $3x10^{-2}$ %. The distribution was truncated at plus and minus 2 standard deviations to avoid cases that are not considered physically possible. The random curves are generated by sampling the log normal distribution, computing the scale factor on the modulus or damping at $3x10^{-2}$ % shear strain, and then applying this factor at all strains. The random variations are reduced at the ends of the strain range to preserve the general shape of the base case curves (Silva, 1992).

3 Summary of Resulting Amplification Factors

3.1 INTRODUCTION

This section presents summary plots of the amplification factors for a subset of the 53 cases. Results of the scaling with V_{S30} and PGA₁₁₇₀ using the cases with randomized soil depths (cases 17, 18, 27, 28, 35, 36, 42, 43) are shown in Section 3.2. Results of the scaling with soil depth for the linear range (e.g., cases with the rock input motion of 0.001g) are shown in Section 3.3. The complete set of results for all cases is attached as an electronic appendix.

3.2 MAGNITUDE SCALING

The period dependence of the amplification for a soil profile with $V_{s30} = 270$ m/sec at the linear range (PGA_{Rock}=0.01g) is shown in Figure 3.1. The amplification for the PR and EPRI models is shown in the top and bottom panels, respectively. In each panel, the amplification from three point-source magnitudes is presented. There is very little difference in the amplification for this low shaking level at periods smaller than 2sec. The increase in amplification for the M5.0 input at periods greater than 2sec is believed to be a numerical artifact of the simulation procedure and therefore results in that range (M=5, T>2sec) are not used for regression of the parametric model.

At higher shaking levels, such as $PGA_{Rock}=1.0g$ (see Figure 3.2), the difference in amplification for the M5.0 simulations is more significant, with a clear resonant peak at T=1 sec. It should be noted here again, that for a M5.0 point source to reach PGA_{Rock} of 1.0g, the source-to-site distance must be much shorter than for the larger point-source magnitudes, leading to such differences.

Figures 3.3 and 3.4 show a similar comparison for a profile with V_{s30} =400 m/sec with PGA_{Rock} = 0.01g and PGA_{Rock}=1.0g, respectively. The trends are similar to those shown for the softer profile but less significant. For example, for the EPRI model with PGA_{Rock}=1.0g, the increase in the peak amplification between M5.0 and M6.0 is about 30% for the V_{s30}=270 m/sec profile but only about 6% for the V_{s30}=400 m/sec profile.

The Magnitude scaling of the amplification for a PR profile with $V_{s30}=270$ m/sec is shown again in the top panel of Figure 3.5, this time as a function of Rock PGA. The increased Magnitude dependence as input shaking increases is clearly seen. The Magnitude scaling on the PGV amplification, however, shown for the same profile in the bottom panel of Figure 3.5, presents different trends. For PGV amplification, the nonlinearity with respect to Rock PGA has opposite trends between the M5.0 simulations and M6.0-M7.0 simulations. This is related with different frequency contents for different point-source magnitude simulations, resulting with PGV for those simulations being associated with a different frequency range. Since nonlinearity is mostly expected to occur from larger magnitude simulations, the regression for the nonlinear model at PGV will be based on M6.0 and M7.0 simulation results only.

For purposes of brevity, all plots in the following sections will present simulation results for a M6.0 point-source only; trends for the M5.0 and M7.0 are similar other than the differences that were mentioned above.



Figure 3.1 Magnitude dependence for a PR, V_{s30} =270 m/sec profile (top) and an EPRI, V_{s30} =270 m/sec profile (bottom) with PGA_{Rock}=0.01g.



Figure 3.2. Magnitude dependence for a PR, V_{S30} =270 m/sec profile (top) and an EPRI, V_{S30} =270 m/sec profile (bottom) with PGA_{Rock}=1.0g.



Figure 3.3 Magnitude dependence for a PR, V_{s30} =400 m/sec profile (top) and an EPRI, V_{s30} =270 m/sec profile (bottom) with PGA_{Rock}=0.01g.



Figure 3.4 Magnitude dependence for a PR, V_{s30} =400 m/sec profile (top) and an EPRI, V_{s30} =270 m/sec profile (bottom) with PGA_{Rock}=1.0g.

3.3 SCALING FOR THE RANDOMIZED SOIL DEPTHS

The period dependence of the amplification in the linear range (PGA₁₁₇₀=0.01g) for the twelve V_{s30} profiles are shown in Figures 3.5 and 3.6. Overall, there is an increase in the amplification as the V_{s30} is reduced.

The nonlinearity in the amplification for a range of V_{S30} profiles and soil models (BM, IV, PR and EPRI) is shown in Figures 3.7 to 3.12 for spectral periods of 0.01, 0.1, 0.2, 0.5, 1.0, and 2.0 sec. The strongest nonlinear effects are seen for T=0.1 sec (Figure 3.8). For spectral

periods of 1.0 sec or longer and soil profiles with V_{S30} of 400 or greater, the amplification is nearly linear. However, in the softer profiles there are two competing trends. For example, for T=2sec the amplification is decreasing with increase in PGA_{Rock} for soil profiles with V_{s30} of 160-190 while it is increasing for profiles with V_{s30} of 270-400 due to the effects of period elongation at these higher levels of shaking.

The nonlinearity of the four soil models can also be compared in Figures 3.7 to 3.12. Note that each V_{s30} value is plotted with the same color on the upper and lower panels. These figures show that on average, the EPRI model is more nonlinear than the PR model. Also, the BM model is much more nonlinear than the IV model. The nonlinearity for the IV model (V_{s30} =190 m/sec) is similar to the nonlinearity for the EPRI model for V_{s30} =270 m/sec.

The standard deviation of the linear amplification $PGA_{Rock}=0.01g$ is shown in Figures 3.13 and 3.14 for the PR soil profiles and the non-PR soil profiles (BM, IV, EPRI), respectively. For the smaller PGA_{Rock} values, the standard deviations of the amplification generally increase with period and decrease with increasing V_{S30} . For profiles with $V_{S30}=270$ m/sec or higher, the standard deviations of the linear amplification are generally between 0.15 and 0.3 natural log units. The PGA_{Rock} dependence of the standard deviations for all soil profiles are shown in Figures 3.15 to 3.17 for spectral periods of 0.01 sec, 0.2 sec and 1.0 sec, respectively. The increase in the standard deviation as a function of the PGA_{Rock} is due to the variability of the nonlinear properties that were randomized (G/G_{max} and damping).


Figure 3.5. Magnitude dependence of the amplification at T=1 sec (top) and at PGV (bottom) for a PR, V_{S30} =270 m/sec profile.



Figure 3.6. Period dependence of the linear amplification for six Vs profiles with PR and depths averaged over 30-1000 ft.



Figure 3.7 Period dependence of the linear amplification for six Vs profiles with BM, IV, and EPRI, and soil depths averaged over 30–1000 ft.



Figure 3.8 PGA_{Rock} dependence of the soil amplification for T=0.01 sec for all soil profiles with a randomized soil depth.



Figure 3.9 PGA_{Rock} dependence of the soil amplification for T=0.1 sec for all soil profiles with a randomized soil depth.



Figure 3.10 PGA_{Rock} dependence of the soil amplification for T=0.2 sec for all soil profiles with a randomized soil depth.



Figure 3.11 PGA_{Rock} dependence of the soil amplification for T=0.5 sec for all soil profiles with a randomized soil depth.



Figure 3.12 PGA_{Rock} dependence of the soil amplification for T=1.0 sec for all soil profiles with a randomized soil depth.



Figure 3.13. Nonlinear amplification for T=2 sec spectral acceleration for the EPRI and PR models using the soil depth averaged over 30-1000 ft.



Figure 3.14 Standard deviations of the linear amplification (PGA_{Rock}=0.01g) for PR with randomized soil depth.



Figure 3.15 Standard deviations of the linear amplification (PGA_{Rock}=0.01g) for BM, IV, and EPRI with randomized soil depth.



Figure 3.16 PGA_{Rock} dependence of the standard deviations for T=0.01 sec for all soil profiles with a randomized soil depth.



Figure 3.17 PGA_{Rock} dependence of the standard deviations for T=0.2 sec for all soil profiles with a randomized soil depth.

3.4 SOIL DEPTH SCALING

The 1-D simulation results can also be used to estimate the effect of the soil depth on the amplification. Examples of the effect of the soil depth on the amplification are shown in Figures 3.18 to 3.20, for soil profiles with V_{S30} ranging from 160 m/sec to 760 m/sec. Similar plots show the effect of depth on the standard deviation of the linear amplification in Figures 3.21 to 3.23. The soil depth is parameterized by the depth to the layer with $V_S=1.0$ km/sec. These figures are for the linear site response (PGA_{Rock}=0.01g); therefore, they are independent of the nonlinear properties.

The soil depth scaling is shown in Figure 3.24 for soil profiles with V_{S30} values of 190 m/sec, 270 m/sec, 400 m/sec, and 560 m/sec. The soil depth effects are strongest for the softer soil profiles (lower V_{S30} values). For the softer profiles, there is strong scaling for soil depths from 8 to 300 m, but the scaling becomes weak for soil depths greater than 300 m, mostly due to the limitations of modeling deep soil profiles with a 1D soil column.



Figure 3.18 PGA_{Rock} dependence of the standard deviations for T=1.0 sec for all soil profiles with a randomized soil depth.



Figure 3.19 Linear amplification for a range of soil depths for the V_{S30} =160 m/sec profile (top) and the V_{S30} =190 m/sec profile (bottom).



Figure 3.20 Linear amplification for a range of soil depths for the V_{S30}=270 m/sec profile (top) and the V_{S30}=400 m/sec profile (bottom).



Figure 3.21 Linear amplification for a range of soil depths for the V_{S30}=560 m/sec profile (top) and the V_{S30}=760 m/sec profile (bottom).



Figure 3.22 Standard deviations of the linear amplification (PGA_{Rock}=0.01g) for a range of soil depths for the V_{S30} =160 m/sec profile (top) and the V_{S30} =190 m/sec profile (bottom).



Figure 3.23 Standard deviations of the linear amplification (PGA_{Rock}=0.01g) for a range of soil depths for the V_{S30} =270 m/sec profile (top) and the V_{S30} =400 m/sec profile (bottom).



Figure 3.24 Standard deviations of the linear amplification (PGA_{Rock}=0.01g) for a range of soil depths for the V_{S30} =560 m/sec profile (top) and the V_{S30} =760 m/sec profile (bottom).



Figure 3.25 Soil depth dependence of the linear amplification.

4 Parametric Model for Nonlinear Site Amplification

4.1 INTRODUCTION

A subset of the simulation cases described above was used for the development of parametric models for the nonlinear response of soil sites to strong ground motions. The proposed models are updates to the models developed by Walling et al. (2008) in that they are based on an extended simulation dataset and use the same functional form.

The models developed by Walling et al. (2008) were revisited herein for two main reasons: (a) the simulation database on which the models were developed has been updated and extended, with additional magnitudes and soil profiles, and (b) two alternatives for the input shaking parameter are explored – the Peak Ground Acceleration (PGA) and the spectral acceleration for each period, (Sa(T)).

4.2 MODEL DEVELOPMENT

Four resulting models are presented below: for each of the soil models - Peninsular Range and EPRI - there is a one model based on PGA as input and one based on Sa(T) as input. All models are based on the randomized-depth soil profiles only. A summary of the simulation cases used for model development is presented in Table 4.1.

V _{s30} (m/sec)	Depth to top of rock (V _{s30} =1 km/sec)	Material model used for nonlinear properties					
190	9-305 m	PR					
270	9-305 m	PR, EPRI					
400	9-305 m	PR, EPRI					
560	9-305 m	PR, EPRI					
760	79 m	PR, EPRI					
900	79 m	PR					

Table 4.1List of simulation scenarios that were selected for model
development.

4.2.1 Functional Form

The functional form is identical to that used in Walling et al. (2008). It can be written as the sum of a linear term and a nonlinear term:

$$\ln(Amp) = f_L(V_{s30}) + f_{NL}(GM_{Rock}, V_{s30})$$
(4.1)

The linear term is a function of Vs30 only, whereas the nonlinear term is a function of Vs30 and a measure of the shaking intensity on rock. The ground motion intensity was defined in the Walling2008 model in term of PGA_{Rock}. Here we present two alternatives – for each soil model (PR and EPRI), we present one model in term of PGA_{Rock} and one in term of Sa(T)_{Rock}. Finally, the resulting functional forms for the two alternative models are given in Equations (4.2) and (4.3) for the PGA-based and the Sa(T)-based models, respectively:

$$\ln(Amp) = \begin{cases} a \ln\left(\frac{V_{S30}^*}{V_{Lin}}\right) - b \ln(PGA_{Rock} + c) \\ + b \ln\left(PGA_{Rock} + c\left(\frac{V_{S30}^*}{V_{Lin}}\right)^n\right) + d \quad for \, V_{s30} < V_{Lin} \\ (a + bn)\ln\left(\frac{V_{S30}^*}{V_{Lin}}\right) + d \quad for \, V_{s30} \ge V_{Lin} \end{cases}$$

$$\ln(Amp) = \begin{cases} a \ln\left(\frac{V_{S30}^*}{V_{Lin}}\right) - b \ln(Sa_{Rock}(T) + c) \\ + b \ln\left(Sa_{Rock}(T) + c\left(\frac{V_{S30}^*}{V_{Lin}}\right)^n\right) + d \quad for \, V_{s30} < V_{Lin} \\ (a + bn)\ln\left(\frac{V_{S30}^*}{V_{Lin}}\right) + d \quad for \, V_{s30} \ge V_{Lin} \end{cases}$$

$$(4.3)$$

where:

$$V_{S30}^* = \begin{cases} V_{S30} & \text{for } V_{S30} < V_1 \\ V_1 & \text{for } V_{S30} \ge V_1 \end{cases}$$
(4.4)

and V_1 corresponds to the $V_{s_{30}}$ above which the soil amplification no longer scales linearly with respect to changes in $V_{s_{30}}$ (Abrahamson and Silva, 2008).

Further explanation on the evolution of the form for the nonlinear term is provided in Walling et al. (2008).

4.2.2 Data Selection

For the models presented below, only simulations with randomized depth (30–1000 ft) are used. The simulations with smaller depth bins are used to constrain the depth term within the GMPE (e.g., Abrahamson and Silva 2008) but that part is not discussed in this report.

A Magnitude-Period constrain was placed on the dataset used for the model regression, due to an increased amplification at long periods for the smaller magnitudes (e.g., Figures 3.1 to 3.4), which is believed to be a numerical artifact of the simulation procedure. The simulation results used for model regression include the Magnitude 7.0 simulations within the available

period range 0.01 < T < 10 sec, Magnitude 6.0 simulations at 0.01 < T < 5 sec and Magnitude 5.0 simulations at 0.01 < T < 2 sec. The PGV model regression is performed on simulations from Magnitudes 6.0 and 7.0 only, as discussed in Section 3.2.

4.2.3 Model Parameters

Due to the nonlinearity of the functional form as shown in figure (4.2) and (4.3), the three parameters - b, n, and c - which appear in the nonlinear term are highly correlated and hence do not all need to be period-dependent. Hence, the parameters n and c were fixed along all periods and the parameter b was regressed as a period-dependent parameter. Fixing of n and c was done based on regressions of smaller subsets at short periods, where the nonlinearity is greatest. In the second step, the period-dependent reference shear-wave velocity, V_{Lin} was regressed and smoothed. V_{Lin} is intentionally kept identical for the two models of each soil group (i.e. the two PR models and the two EPRI models), since it represents a physical measure which should not depend on the input shaking. In the third step, the period-dependent nonlinear coefficient, *b*, was regressed and smoothed.

Smoothing of the period-dependent parameters was done by fitting a 7th order polynomial to the regressed values with some constraints at the low and high ends of the period range, following Equation (4.5).

$$x = \begin{cases} \beta_2 & T \ge T_2 \\ \alpha_0 + \sum_{i=1}^7 \alpha_i \left[ln \left(\frac{T}{T_0} \right) \right]^i & T_1 < T < T_2 \\ \beta_1 & T \le T_1 \end{cases}$$
(4.5)

The smoothed parameter x in Equation (4.5) is either $ln(V_{Lin})$ or b and T is the spectral period. The four pairs of the smoothed parameters (V_{lin} and b for each of the four models presented in this report) are compared with the smoothed parameters from Walling et al. (2008) in Figure 4.1. Note that while the parameter b was constrained in Walling et al. (2008) to be negative (or \leq 0.15), it was allowed to be positive in the models below and was as high as 3.95 for the PR-Sa model. The positive b values allow an increase in amplification for increased levels of shaking, which can be related to period elongation, as will be shown below.

The list of parameters needed for reconstructing the nonlinear terms (following equation 5) are presented in Table 4.2. The smoothed values for V_{lin} and b are listed in Appendix A for the 111 periods that are in the NGA FlatFile. The parameters which appear in the linear term - a and d - are left unsmoothed, since they are correlated with other terms in the GMPE and should be regressed for each individual GMPE in which these site amplification models are applied.



Figure 4.1 Period-dependence of the smoothed model parameters. V_{Lin} (top) and *b* (bottom).

	PR			EPRI		
		PGA	Sa		PGA	Sa
n		1.5	1.5		1.5	1.5
	с	1.4	2.4 (*100 for PGV)		2.0	3.0 (*100 for PGV)
	V _{LIN}	b	b	V _{LIN}	b	b
PGV	332.00	-1.5140	-2.0200	728.00	0.5850	0.6025
T_{θ}	0.010	0.020	0.012	0.014	0.010	0.02
T_{I}	0.015	0.020	0.018	0.018	0.022	0.018
T_2	0.550	9.000	5.500	0.460	1.820	7.000
a_0	6.5300	-1.2500	-1.6400	7.1360	-0.9039	-0.9241
a_1	-0.2000	0.2780	0.9474	-0.6500	1.1276	0.3081
a_2	0.2400	-1.3430	-2.0673	1.7860	-3.5267	0.2166
a_3	0.0940	2.4810	2.2630	-1.0370	4.4341	-0.5068
a_4	-0.0170	-1.8690	-1.0634	0.1237	-2.5880	0.1586
a_5	-0.0529	0.6040	0.2097	0.0421	0.7361	0.0006
a_6	0.0191	-0.0862	-0.0155	-0.0117	-0.0993	-0.0047
a_7	-0.0018	0.0045	0.0002	0.0008	0.0051	0.0004
β_1	6.493	-1.250	-1.470	7.068	-0.833	-0.960
β_2	5.805	0.360	3.950	6.590	0.600	2.100

Table 4.2List of coefficients needed to reconstruct the smoothed model
parameters.

4.3 SUMMARY OF RESULTS

The dependence of the amplification level on shaking intensity [in terms of PGA or Sa(f)] as well as on the soil profile (in terms of $V_{s_{30}}$) is presented in Figures 4.2 through 4.5 for the four different models, for periods 0.01, 0.2, 1, 2, 3, and 5 sec. The simulation results are represented by open symbols while the parametric model is shown by the solid line for each corresponding V_{s30} . All four models can be seen to capture the general response of the simulations, that is – increased amplification as V_{s30} decreases, nonlinearity at short periods, and period elongation (increased amplification for increased input motion) at long periods. The biggest discrepancy can be seen at T=1 sec for all four models. At that period, there are several aspects of the soil response that cannot be captured by the current functional form. For example, the two softest profiles for the PR model (V_{s30} =190 and V_{s30} =270) have opposite trends with shaking intensity, the amplification for low shaking intensity does not scale linearly with V_{s30}, and some of the profiles change their nonlinearity from positive to negative with shaking input. These discrepancies are also consistent with Figures 3.2 and 3.5, which suggest that at T=1 sec there is both resonance at the soft profiles and the greatest magnitude dependence. The model's agreement with simulation results at other periods seems to be much better for a range of V_{s30} and shaking intensity values.



Figure 4.2 Model versus simulation results for the PR-PGA model for six representative periods.



Figure 4.3 Model versus imulation results for the PR-Sa model for six representative periods.



Figure 4.4 Model versus simulation results for the EPRI-PGA model for six representative periods.



Figure 4.5 Model versus imulation results for the EPRI-Sa model for six representative periods.

The nonlinearity with respect to shaking intensity can be represented by the slope of the lines in Figures 4.2 through 4.5. This term is presented in Figure 4.6 for a profile with $V_{s30}=270$ m/sec, showing the slopes between reference PGA of 0.1g to 1g. The trends are similar for all four models and are also quite consistent with nonlinearity computed from the Walling(2008) models. The nonlinearity for the EPRI models is generally greater than that of the PR models, with the nonlinearity peaking at about T=0.2s for all models (as seen in the top panel, slope from 0.1g to 1g).

A comparison between the resulting spectral values on soil vs. the input motion on rock is presented in Figure 4.7 for the PR-Sa model, for the same six periods as shown in Figures 4.2 through 4.5. The expected spectral acceleration of the soil, given the input spectral acceleration on rock and the corresponding model amplification can be computed as:

$$Sa_{Soil}(f) = Sa_{Rock}(f) \cdot Amp(f)$$
(4.6)

It can be seen that at low rock shaking levels, all soil profiles amplify the response (all lines are above the 1:1 line). As shaking levels increase, the softer profiles de-amplify the response at shorter periods (lines cross below the 1:1 line). At the longer periods (above \sim 1 sec) there is no de-amplification even for high levels of shaking.



Figure 4.6 Period-dependence of the combined nonlinear term, computed as the slope of amplification from 0.1g to 1g.



Figure 4.7 Sa on Soil vs. Sa on Rock, for the PR-Sa model.

The spectral shape of the rock and corresponding soil motions are presented in Figure 4.8 for four increasing shaking intensities – ranging from PGA_{1100} of 0.05g to 1g. The rock motions are shown in solid lines and the corresponding soil motions are in dashed lines. It can be seen that as shaking increases, there is more nonlinearity, and hence the soil softens and the peak response is shifted towards longer periods. This can also explain the upward trend of the amplification curve at T>2 sec in Figures 4.3 through 4.6.



PR, Sa model, Vs₃₀ 270

Figure 4.8 Spectral acceleration (top) and Normalized spectral acceleration (bottom) vs. Period, for the PR-Sa model with Vs30=270 m/sec and under four different input shaking intensities.



Figure 4.9 Normalized spectral acceleration vs. Period for a $Vs_{30}=270$ m/sec profile with PGA₁₁₀₀=0.1g (left) and PGA₁₁₀₀=0.5g (right).

The four amplification models are compared in Figure 4.9 in terms of their resulting soil spectra for $Vs_{30}=270$ m/sec with PGA₁₁₀₀=0.1g (left) and PGA₁₁₀₀=0.5g (right). It can be seen that while all four models result in a largely similar soil spectra, there is a greater difference between the two nonlinear material models (i.e. PR vs. EPRI) rather than between the two forms of the input motion [i.e. PGA vs. Sa(T)].

The predictive power of the functional form can be tested by comparing the sigma of the model for different profiles and different functional forms. Since each combination of input shaking value and soil profile was simulated 30 times with randomization of the soil profile and material properties, the parametric model is regressed on the mean amplification of each of those sets. The standard deviation of the fit could therefore either include or not include the standard deviation of each set. A comparison between the two options is presented in Figure 4.10, which shows the standard deviation of the fit for the PR-Sa model, for each of the different soil profiles. On the left, σ_{res} is the standard deviation of the residuals between the simulation means and the parametric model. On the right, the total standard deviation includes the average standard deviation of the simulations for each of the profiles considered, computed as following:

$$\sigma_{res-total} = \sqrt{\overline{\sigma_{simulation}}^2 + {\sigma_{res}}^2} \tag{4.7}$$

where $\overline{\sigma_{simulation}}$ is the mean of the standard deviations for each set of 30 realizations performed for each profile and each shaking intensity (see Figures 3.14 through 3.18) and σ_{res} is the misfit between the mean of that set and the parametric model. In general, Figure 4.10 shows that σ_{res} increases as V_{s30} decreases. The average σ_{res} , representing all soil profiles, is shown in black and can be seen to be mostly affected by the softer profiles with the larger σ_{res} while the average $\sigma_{res-total}$ is pulled down (relatively) by the lower $\sigma_{simulation}$ of the stiffer profiles.

The two different functional forms (PGA vs. Sa as shaking input parameter) are compared in Figure 4.11 for the Peninsular Range soil model, showing both σ_{res} and $\sigma_{res-total}$ for each of the forms. If one functional form would have fit the simulation results better, we would expect that form to have a lower standard deviation. However, both σ_{res} and $\sigma_{res-total}$ are almost identical for the two forms, suggesting that there is no significant advantage for either of the functional forms. These findings are different than the conclusion of Bazzurro and Cornell (2004), who propagated 78 earthquake records through two soil profiles and explore seven functional forms for the amplification functions. A subset of their findings is also re-drawn in Figure 4.11, comparing the standard deviations of the residuals to two of their functional forms, using PGA and Sa(f) as single model parameters. Bazzurro and Cornell concluded that Sa(T) is the single most helpful parameter for the prediction of the amplification since it had the lowest σ_{res} . One of the main difference between our simulations and Bazzurro and Cornell's is that they use real time histories and by that account for variability in the spectral shape, whereas the simulations in this study were performed with the RVT method, which has a unique spectral shape and hence a strong correlation between PGA and Sa(T). Uncertainties in soil properties are accounted for in both studies, and do not add to the total uncertainty very much, as long as the general profile (depth, Vs30) stays the same.

We conclude that based on the analysis presented herein, there is no statistical preference for either of the forms (PGA or Sa as input-motion parameter), but we recommend using the Sa(T) model, for ease of use in forward applications of the GMPE model.


Figure 4.10 Standard deviation of PR-Sa model, for a range of V_{s30} values. On the left σ_{res} is the standard deviation of the residuals to the fit, while on the right it is the total standard deviation, including the standard deviation of each set of 30 realizations.



Figure 4.11 Standard deviation of the PR models, averaged over all Vs30 profiles. Comparison between the two PR models from the current study and a subset of two corresponding models redrawn from Bazzurro and Cornell (2004).

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Appendix A: Table of Model Parameters

	PR PGA Model		PR Sa Model		EPRI PGA Model		EPRI Sa Model	
n	1.5		1.5		1.5		1.5	
С	1.4		2.4 (*100 for PGV)		2.0		3.0 (*100 for PGV)	
Period (sec)	VLIN	b	VLIN	b	VLIN	b	VLIN	b
PGV	332.00	-1.514	332.00	-2.020	728.00	0.585	728.00	0.603
0.01	660.50	-1.250	660.50	-1.470	1173.80	-0.833	1173.80	-0.960
0.02	683.74	-1.250	683.74	-1.459	1195.76	-0.833	1195.76	-0.924
0.022	698.64	-1.234	698.64	-1.448	1232.97	-0.833	1232.97	-0.893
0.025	723.96	-1.232	723.96	-1.430	1304.69	-0.825	1304.69	-0.850
0.029	760.05	-1.236	760.05	-1.399	1409.62	-0.805	1409.62	-0.803
0.03	769.14	-1.237	769.14	-1.390	1435.59	-0.799	1435.59	-0.793
0.032	787.11	-1.237	787.11	-1.372	1486.02	-0.787	1486.02	-0.776
0.035	813.17	-1.233	813.17	-1.343	1556.26	-0.770	1556.26	-0.757
0.036	821.52	-1.231	821.52	-1.334	1577.95	-0.765	1577.95	-0.752
0.04	852.89	-1.220	852.89	-1.297	1655.02	-0.749	1655.02	-0.739
0.042	867.19	-1.214	867.19	-1.279	1687.55	-0.744	1687.55	-0.736
0.044	880.48	-1.208	880.48	-1.263	1716.06	-0.740	1716.06	-0.735
0.045	886.74	-1.205	886.74	-1.255	1728.85	-0.738	1728.85	-0.735
0.046	892.74	-1.201	892.74	-1.247	1740.67	-0.737	1740.67	-0.735
0.048	903.95	-1.196	903.95	-1.232	1761.51	-0.737	1761.51	-0.737
0.05	914.11	-1.190	914.11	-1.219	1778.77	-0.738	1778.77	-0.740
0.055	934.95	-1.180	934.95	-1.191	1807.65	-0.746	1807.65	-0.753
0.06	949.57	-1.175	949.57	-1.170	1818.90	-0.763	1818.90	-0.772
0.065	958.49	-1.176	958.49	-1.157	1815.95	-0.786	1815.95	-0.794
0.067	960.59	-1.178	960.59	-1.154	1811.51	-0.796	1811.51	-0.803
0.07	962.31	-1.183	962.31	-1.151	1801.90	-0.814	1801.90	-0.818
0.075	961.69	-1.196	961.69	-1.152	1779.42	-0.845	1779.42	-0.844
0.08	957.26	-1.214	957.26	-1.158	1750.74	-0.879	1750.74	-0.872
0.085	949.63	-1.237	949.63	-1.169	1717.67	-0.915	1717.67	-0.899
0.09	939.35	-1.264	939.35	-1.186	1681.64	-0.952	1681.64	-0.926
0.095	926.93	-1.294	926.93	-1.206	1643.80	-0.990	1643.80	-0.953
0.1	912.81	-1.328	912.81	-1.230	1605.04	-1.027	1605.04	-0.980

Table A.1. Model Parameters corresponding to the 111 periods in the NGA FlatFile.

0.11	880.95	-1.404	880.95	-1.287	1527.26	-1.101	1527.26	-1.031
0.12	846.25	-1.487	846.25	-1.353	1451.92	-1.171	1451.92	-1.078
0.13	810.51	-1.574	810.51	-1.426	1380.93	-1.236	1380.93	-1.121
0.133	799.78	-1.601	799.78	-1.450	1360.64	-1.255	1360.64	-1.134
0.14	774.94	-1.664	774.94	-1.505	1315.18	-1.296	1315.18	-1.161
0.15	740.37	-1.755	740.37	-1.587	1254.92	-1.349	1254.92	-1.196
0.16	707.30	-1.846	707.30	-1.671	1200.08	-1.397	1200.08	-1.228
0.17	676.04	-1.935	676.04	-1.756	1150.38	-1.439	1150.38	-1.256
0.18	646.74	-2.022	646.74	-1.842	1105.46	-1.475	1105.46	-1.280
0.19	619.44	-2.107	619.44	-1.927	1064.92	-1.505	1064.92	-1.301
0.2	594.13	-2.188	594.13	-2.012	1028.37	-1.530	1028.37	-1.319
0.22	549.18	-2.342	549.18	-2.177	965.74	-1.567	965.74	-1.347
0.24	511.11	-2.481	511.11	-2.335	914.87	-1.586	914.87	-1.366
0.25	494.37	-2.545	494.37	-2.411	893.14	-1.590	893.14	-1.372
0.26	479.03	-2.606	479.03	-2.485	873.56	-1.591	873.56	-1.376
0.28	452.08	-2.717	452.08	-2.626	840.03	-1.584	840.03	-1.379
0.29	440.28	-2.768	440.28	-2.693	825.73	-1.576	825.73	-1.378
0.3	429.47	-2.815	429.47	-2.757	812.88	-1.566	812.88	-1.376
0.32	410.55	-2.900	410.55	-2.879	791.00	-1.540	791.00	-1.367
0.34	394.74	-2.973	394.74	-2.992	773.51	-1.507	773.51	-1.354
0.35	387.85	-3.005	387.85	-3.045	766.18	-1.488	766.18	-1.346
0.36	381.57	-3.034	381.57	-3.096	759.70	-1.468	759.70	-1.338
0.38	370.63	-3.086	370.63	-3.191	749.01	-1.424	749.01	-1.318
0.4	361.60	-3.128	361.60	-3.278	740.98	-1.376	740.98	-1.295
0.42	354.20	-3.162	354.20	-3.356	735.24	-1.326	735.24	-1.270
0.44	348.19	-3.187	348.19	-3.427	731.48	-1.273	731.48	-1.243
0.45	345.64	-3.197	345.64	-3.460	730.26	-1.246	730.26	-1.229
0.46	343.37	-3.205	343.37	-3.491	727.78	-1.218	727.78	-1.214
0.48	339.57	-3.217	339.57	-3.548	727.78	-1.162	727.78	-1.184
0.5	336.64	-3.222	336.64	-3.599	727.78	-1.106	727.78	-1.153
0.55	331.96	-3.212	331.96	-3.698	727.78	-0.963	727.78	-1.070
0.6	331.96	-3.174	331.96	-3.765	727.78	-0.822	727.78	-0.983
0.65	331.96	-3.114	331.96	-3.802	727.78	-0.685	727.78	-0.895
0.667	331.96	-3.089	331.96	-3.808	727.78	-0.640	727.78	-0.864
0.7	331.96	-3.037	331.96	-3.814	727.78	-0.554	727.78	-0.805
0.75	331.96	-2.946	331.96	-3.804	727.78	-0.431	727.78	-0.716
0.8	331.96	-2.845	331.96	-3.775	727.78	-0.316	727.78	-0.627
0.85	331.96	-2.737	331.96	-3.729	727.78	-0.209	727.78	-0.540
0.9	331.96	-2.622	331.96	-3.670	727.78	-0.111	727.78	-0.454
0.95	331.96	-2.504	331.96	-3.598	727.78	-0.021	727.78	-0.370
1	331.96	-2.383	331.96	-3.515	727.78	0.061	727.78	-0.288
1.1	331.96	-2.137	331.96	-3.323	727.78	0.202	727.78	-0.131
1.2	331.96	-1.892	331.96	-3.104	727.78	0.316	727.78	0.018
1.3	331.96	-1.653	331.96	-2.865	727.78	0.405	727.78	0.157
1.4	331.96	-1.421	331.96	-2.611	727.78	0.473	727.78	0.287

1.5	331.96	-1.199	331.96	-2.348	727.78	0.524	727.78	0.409
1.6	331.96	-0.989	331.96	-2.079	727.78	0.559	727.78	0.523
1.7	331.96	-0.791	331.96	-1.807	727.78	0.583	727.78	0.629
1.8	331.96	-0.605	331.96	-1.534	727.78	0.597	727.78	0.728
1.9	331.96	-0.431	331.96	-1.262	727.78	0.600	727.78	0.821
2	331.96	-0.269	331.96	-0.992	727.78	0.600	727.78	0.907
2.2	331.96	0.020	331.96	-0.466	727.78	0.600	727.78	1.062
2.4	331.96	0.264	331.96	0.038	727.78	0.600	727.78	1.198
2.5	331.96	0.371	331.96	0.281	727.78	0.600	727.78	1.259
2.6	331.96	0.468	331.96	0.516	727.78	0.600	727.78	1.315
2.8	331.96	0.636	331.96	0.963	727.78	0.600	727.78	1.418
3	331.96	0.772	331.96	1.379	727.78	0.600	727.78	1.507
3.2	331.96	0.879	331.96	1.763	727.78	0.600	727.78	1.585
3.4	331.96	0.961	331.96	2.114	727.78	0.600	727.78	1.653
3.5	331.96	0.993	331.96	2.277	727.78	0.600	727.78	1.684
3.6	331.96	1.020	331.96	2.432	727.78	0.600	727.78	1.712
3.8	331.96	1.060	331.96	2.718	727.78	0.600	727.78	1.764
4	331.96	1.084	331.96	2.972	727.78	0.600	727.78	1.809
4.2	331.96	1.093	331.96	3.195	727.78	0.600	727.78	1.848
4.4	331.96	1.089	331.96	3.388	727.78	0.600	727.78	1.883
4.6	331.96	1.076	331.96	3.551	727.78	0.600	727.78	1.913
4.8	331.96	1.053	331.96	3.685	727.78	0.600	727.78	1.939
5	331.96	1.024	331.96	3.792	727.78	0.600	727.78	1.963
5.5	331.96	0.927	331.96	3.950	727.78	0.600	727.78	2.010
6	331.96	0.813	331.96	3.950	727.78	0.600	727.78	2.046
6.5	331.96	0.694	331.96	3.950	727.78	0.600	727.78	2.075
7	331.96	0.584	331.96	3.950	727.78	0.600	727.78	2.100
7.5	331.96	0.490	331.96	3.950	727.78	0.600	727.78	2.100
8	331.96	0.418	331.96	3.950	727.78	0.600	727.78	2.100
8.5	331.96	0.376	331.96	3.950	727.78	0.600	727.78	2.100
9	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
9.5	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
10	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
11	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
12	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
13	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
14	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
15	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100
20	331.96	0.360	331.96	3.950	727.78	0.600	727.78	2.100

Appendix B: Table of Amplification Factors for Two Hard Rock Profiles

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Appendix C: Table of Amplification Factors for the Soil Profiles by Depth Bins

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Appendix D: Table of Amplification Factors for the Soil Profiles with Randomized Depth

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